

Article

Auriferous quartz veining due to decompressional cooling and CO₂ content variations revealed by quartz solubility, SEM-CL and fluid inclusion analyses (the Linglong goldfield)

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Abstract: Quartz is the most common gangue mineral in hydrothermal veins. Coupled with capacities of hosting fluid inclusions and recording varieties of microtextures, its solubility behavior may provide unparalleled insights into hydrothermal processes. In this study, the Linglong goldfield is targeted to investigate gold-producing quartz veining process. Scanning electron microscope (SEM)-cathodoluminescence (CL) imaging uncovered three episodes of quartz deposition, intervened by an episode of quartz dissolution. Based on newly-developed quartz solubility diagrams and CL-aided fluid inclusion microthermometry, it is proposed that the earliest quartz (Qz1) was formed by a combined effect of CO₂ increase and decompressional cooling, leading to the formation of the early thick gold-barren veins (V1); the second generation of quartz (Qz2a) was formed by the same fluids that may have been diluted and cooled by meteoric water, leading to much reduced quantity of quartz and deposition of pyrite and gold; and the third generation of quartz (Qz2b) was deposited along with polymetallic sulfides due to fluid cooling following a quartz dissolution event likely induced by cooling in retrograde solubility region and/or CO₂ decrease. This research may help understand gold formation in orogenic intrusion related deposits, and points to imperative of CL-based in-situ analysis for future studies.

Keywords: Mineral solubility; hydrothermal veins; Jiaodong gold; fluid overprint

1. Introduction

Quartz is the most important silica mineral and widely occurs in igneous, metamorphic, sedimentary and hydrothermal rocks of the Earth crust. Because of its abundance and properties, quartz and silica minerals find many applications in industry and scientific researches [1]. For instance, its luminescence and geochemical properties and host fluid/mineral inclusions have provided critical knowledge for reconstructing geological processes [2-5]. In hydrothermal ore deposits, quartz is the most common gangue mineral and may intergrow with metals (such as native gold) and sulfides containing metals such as copper, tin, and tungsten [6]. The veining processes of quartz and associated minerals in porphyry Cu deposits have been intensively studied and well understood with the analytical aid of cathodoluminescence (CL) petrography [7-9], solubility analysis [10,11], fluid inclusion [12,13], trace elements [14] and oxygen isotopic composition [15,16]. These pieces of information ultimately shed lights on where cupriferous fluids are sourced and how these fluids precipitate metal sulfides and quartz. By contrast, quartz solubility analysis have been

much less employed in the studies of orogenic intrusion-related gold deposits, leading to less knowledge of the veining processes that are responsible for gold deposition. Several previous studies were based on quartz solubility in pure water rather than H_2O - NaCl - CO_2 ternary system that is a better approximation of orogenic gold fluid [17]. Recently, Monecke et al. [18] constructed solubility diagrams in the H_2O - NaCl - CO_2 ternary system but data are only available for the single-phase region. This lack of data for two-phase field poses a challenge for thoroughly understanding vein formation processes in orogenic intrusion-related setting, since fluid immiscibility is quite common and is considered a critical factor for gold genesis [19].

The Linglong gold deposit of the Jiaodong gold province is targeted in this study to fill this void of knowledge. This deposit is selected because (1) it is a vein-type gold deposits, comprising large quartz veins that host gold mineralizations; and (2) its geology, geochemistry and isotopic characteristics have been well documented [20-27]. Based on a methodological combination of CL petrography, fluid inclusion and newly-developed quartz solubility diagrams that cover both the single-phase and two-phase field, this study elucidates veining processes in this deposit, which have implications for understanding quartz deposition and dissolution processes in a wide range of settings.

2. The Linglong gold deposit

2.1. General geology

The Jiaodong gold province (JGP) has been the subject of detailed studies for many decades [28-31] and its geological and metallogenic backgrounds have been synthesized and updated by many past and recent works [32-44]. Inasmuch, only a brief introduction closely relevant to this study is provided. The region of the JGP is a peninsula located in the northeastern China (Figure 1a inset). Tectonically, it straddles southeastern part of the Precambrian North China Craton (NCC) in the north and the Triassic Sulu ultra - high pressure orogenic belt (UHP) in the south; and to the west, it is bounded by the north-northeast trending Tan-Lu fault zone [45] (Figure 1a). This region is considered to have been involved in the Pacific tectonic regime since early Mesozoic [46,47]. Rocks exposed include basement rocks metamorphosed at high amphibolite to granulite facies in the early Proterozoic (1.80 – 1.95 Ga, [48]), Triassic UHP rocks [49] and voluminous Jurassic to Cretaceous volcanic-sedimentary rocks [50], granitoids [51] and mafic to intermediate dykes [52] (Figure 1a).

Gold deposits in Jiaodong are structurally controlled within a set of NE- and NNE- trending fault systems [53], forming three clusters, from west to east, namely the Zhaoyuan - Laizhou, Penglai - Qixia and Muping - Rushan belts [54]. Over 95% of the gold resources are hosted in the Cretaceous granitoids [55], with only a few small-sized deposits being hosted in the metamorphic basement [56,57]. Gold mineralization occurs either as auriferous quartz veins (Linglong-type) or as disseminated/stockwork ores rocks (Jiaojia-type) [29] (Figure 1a). The difference in mineralization styles, for instance in the Linglong goldfield, may have been caused by difference in local stress field and the related physico-chemical variations [26].

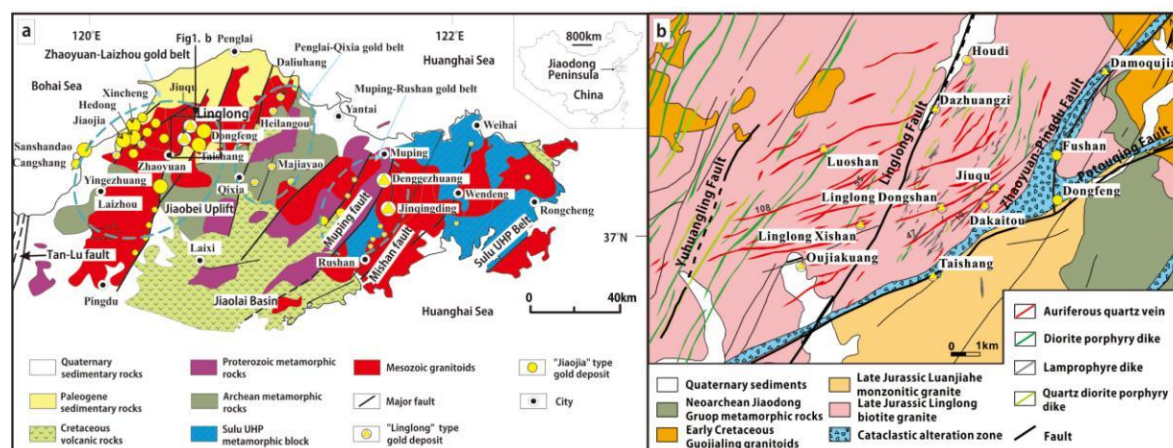


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Figure 1. Geological maps of the Jiaodong gold province and the Linglong goldfield. (a) Geology and spatial distribution of gold deposits in the Jiaodong gold province, which is located in eastern China as a peninsula (inset). Filled circles represent disseminated ore deposits; circles filled with a triangle represent vein-type deposits. Symbol sizes are proportional to gold reserves. (b) A geological map of the Linglong goldfield with locations of known ore deposits.

The Linglong goldfield, the largest gold producing area in Jiaodong, is located at the northern tip of the Zhaoyuan - Pingdu sub-belt (ZPB), which commonly hosts disseminated gold ores (Figure 1b). This goldfield is characterized by the presence of hundreds of gold-producing quartz veins, occurring in steeply dipping, NE - NNE trending extensional cracks in the hanging wall of ZPB. These veins are currently under exploration and production by several mines such as Dongshan, the Xishan, the Dakaitou, and the Jiuqu. In the eastern part, the ZPB is joint by the Potouqing fault, the conjunction of which hosts disseminated ores [26]. Most of the gold deposits are usually hosted in the Linglong biotite granite, with deep ores, for instance at Jiuqu, being hosted in both the Linglong biotite granite and Guojialing granodiorite [26]. The Linglong fault crosscut the ZPB and its subordinate faults controls the emplacement of intermediate to mafic dykes.

Over two hundred veins are exposed to the surface in the Dongshan and Xishan mines, among which roughly 30 veins have economic values, such as the veins #108, #55, #53, #52, #51, #48, #47, #10, #9 (Figure 1b). In general, the veins mostly strike NE between 35° to 70° and dipping NW, whereas may change to SE-dipping at shallows and NW-dipping at depth near the Potouqing Fault. These veins often reach thousands of meters in length and have variable widths of one to tens of meters. Ore grades vary from a few grams to tens of grams per ton up to hundreds of grams per ton, where visible from naked eyes.

2.2. Alterations, veins and mineralization

Hydrothermal alteration, veins and mineralization characteristics of disseminated ore bodies have been detailed by [26] and [27]. This section primarily focuses on vein-type mineralizations and the associated alterations, haloes.

Hydrothermal alterations observed at Linglong include potassic (mainly K-feldsparization and biotitization), sericitic, silicic, pyritic alterations and carbonatization. The potassic alteration generally developed as the outer part of the alteration zones in the Linglong granite, and often overprinted by sericitic and silicic alterations (Figure 2a). Potassic alteration is characterized by conversion of magmatic plagioclase and biotite into secondary K-feldspar and biotite (Figure 2b). This alteration resulted in dramatic increases in K-feldspar and biotite in the host Linglong granite. Sericitic alteration commonly shows light green color, which converts feldspar into fine-grained sericite. Sericitic alteration is marked by its green to white colors and overprints the potassic alterations (Figure 2c). Silicification can occur along with other alterations, but also individually. They were characterized by replacement of previous minerals by quartz (Figure 2d). Its intensity increases towards the thick quartz veins. These alterations commonly juxtapose to form symmetrical zones according to the central thick quartz veins (Figure 2a), which, towards fresh granite, comprise pyrite - sericite - quartz, sericite - quartz, and K-feldspar - quartz. This alteration types and spatial distribution are comparable to that of disseminated gold deposits such as the Sanshandao [58] and Dongfeng deposits [26], but differ with much small thickness.

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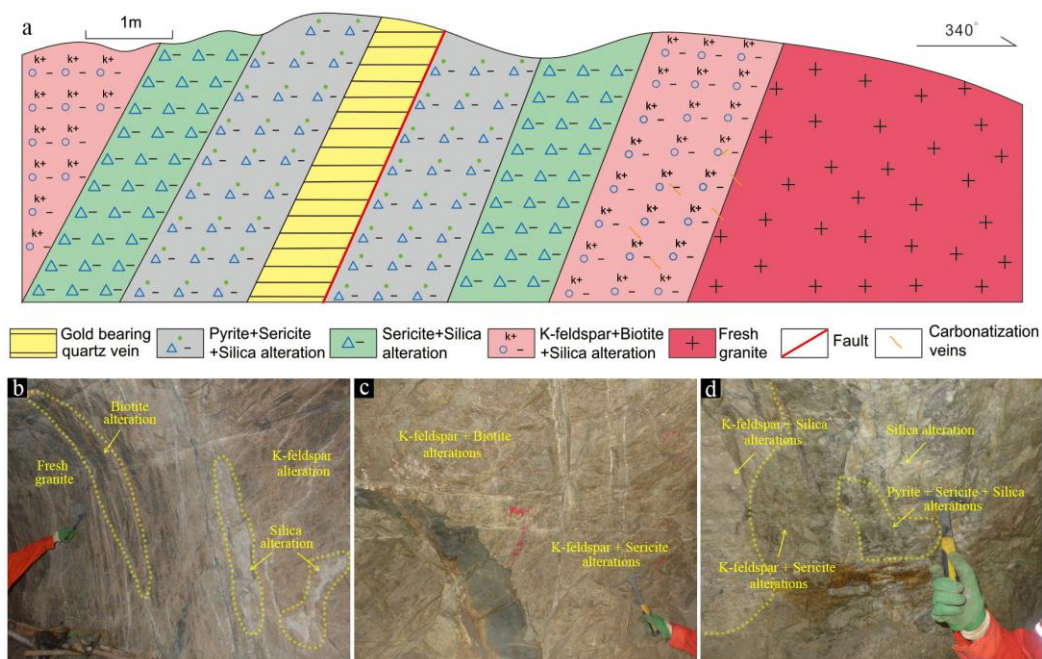


Figure 2. A sketch of alteration zoning (a) and underground-tunnel images of the alterations at the Linglong goldfield (b-d). (b) Pink to flesh pink potassic alterations developed in the biotite granite and were overprinted by the silicic alterations. (c) Sericitization overprinted on the K-feldspar alterations. (d) The silicification as the quartz replaced the potassic and sericite alterations with local pyritic alterations.

Based on field (underground tunnels and outcrops) and hand specimen observations and crosscutting relations, four generations of vein have been identified, from early to late, including quartz $\pm$ pyrite (V1), gold + quartz + pyrite (V2), gold + quartz + base metal sulfides (V3) and quartz + calcite $\pm$ pyrite (V4) at the Linglong goldfield (Figure 3a). V1 veins often cracked and developed with potassic halos (Figure 3b); V2 veins usually thin with several centimeters and take on liner distributions, pyrite $\pm$ quartz $\pm$ sericite alterations were seen appeared around V2 (Figure 3c); The cutting relationship showed that V1 were cut by V2 (Figure 3d).

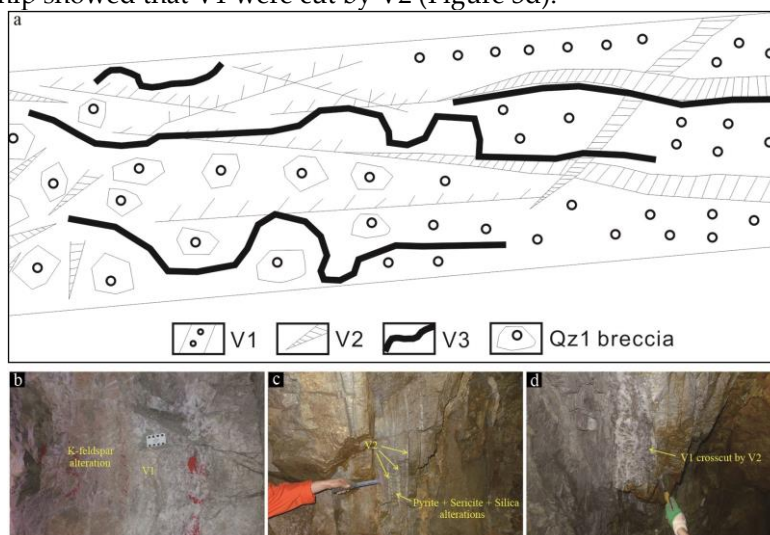


Figure 3. A sketch of vein types and their crosscutting relation (a) and underground-tunnel images of the vein generations of the Linglong goldfield. (b) Cracked V1 developed with K-feldspar alterations. (c) V2 overprinted by the pyrite + sericite + Silica alterations. (d) V1 crosscut by V2.

In V1, quartz accounts > 95 vol.% of vein minerals, with minor amounts of coarse-grained anhedral or subhedral pyrite. Little gold was precipitated with this vein type. The quartz grains are

subhedral to euhedral in morphology and display milky to white color (Figure 4a). Pyrite grains are commonly fractured. V2 veins consist of about 20 - 30 vol.% quartz, 70 vol.% pyrite. Gold mainly occurs as native gold in pyrite fractures. Quartz in this vein type show white-gray color and poor transparency. Pyrite grains show euhedral cubes and semi-euhedral aggregates and often cement or crosscut the V1 (Figure 4b). V3 veins are characterized by the appearance of base-metal sulfides (pyrite, chalcopyrite, galena, sphalerite, minor pyrrhotite and tetrahedrite). Metal sulfides occurs with gold and few quartz (about 5 vol.%) (Figure 4c). Quartz grains of this vein type usually dark to gray in color and their relative volume is significantly decreased. Pyrite and other metal sulfides are fine-grained semi-euhedral and anhedral aggregates. V2 and V3 veins are the major sources of gold. V4 veins could be observed in the mafic dykes which consist of carbonatizations (Figure 4d). Pyrite and gold are sporadically seen.

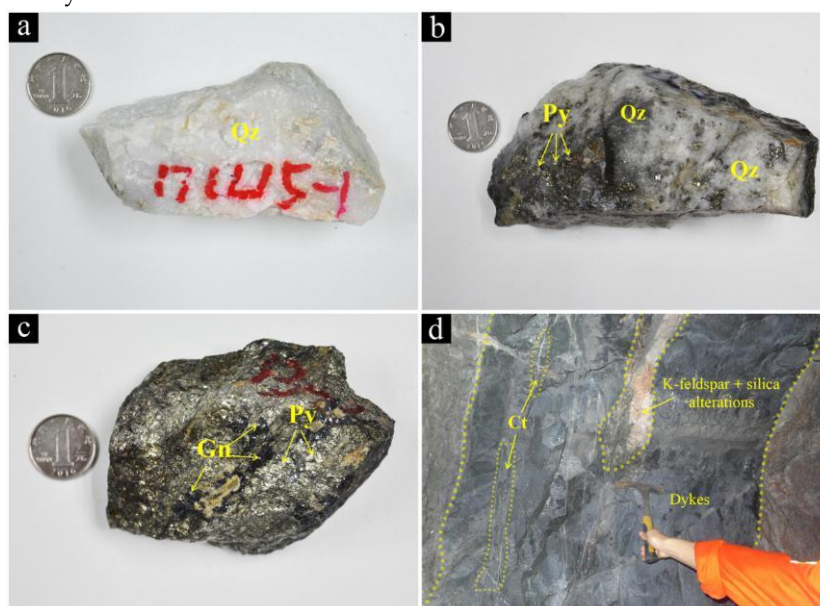


Figure 4. The ores and the veins in different stages of mineralization at the Linglong goldfield. (a) A hand specimen of V1 with milky white quartz; (b) A hand specimen of V2 with quartz in white-gray color and poor transparency; (c) A hand specimen of V3 appearance of base-metal sulfides with golden color. (d) A tunnel picture of V4 in dykes and overprinted by potassic and silicic alterations. Qz: quartz, Py: pyrite, Ct: calcite, Gn: galena.

3. Materials and Methods

3.1. Samples and analytical protocols

Twenty-nine vein and rock samples were selected from underground tunnels of the Xishan (14), Jiuqu (15) mines in Linglong goldfield. Thin and thick sections were made for petrographic observations by using microscope, among of which eleven sections were carefully studied by scanning electron microscope (SEM) equipped with a cathodoluminescence (CL) detector. Eight quartz samples were determined by using cooling/heating stages and laser Raman spectrometer for fluid pressure, temperature and composition. A compilation of petrography of these samples are given in the Table 1.

Table 1. Samples and their vein generation and mineral assemblage of this study.

Sample No.	Vein generations	Mineral assemblage
10LL23	V1	pyrite
17LL15	V1	pyrite
17LL21	V1	pyrite
17LL05-01	V1 + V2	pyrite
17LL05-02	V1 + V2	pyrite

17LL01	V2		pyrite
10LL04	V2	pyrite	
17LL07	V2		pyrite
17LL09	V2		pyrite
17LL10-01	V2 + V3		pyrite, siderite
17LL10-02	V2 + V3		pyrite
17LL10-03	V2 + V3		pyrite
17LL10-04	V2 + V3		pyrite, chalcopyrite, galena
17LL13-01	V3	pyrite, siderite, sphalerite, chalcopyrite, galena, pyrrhotite, Argentum	
17LL13-02	V3	pyrite, siderite, sphalerite, galena, chalcopyrite, tetrahedrite, pyrrhotite, Aurum, Argentum	
12LL10-01	V4		pyrite, sphalerite, galena
12LL10-02	V4		pyrite, sphalerite
12LL11-01	V4		pyrite
12LL11-02	V4		pyrite

Microthermometric measurements and Raman spectroscopy analyses on the fluid inclusions were carried out using a Linkam THMSG 600 heating-freezing stage combined with a Zeiss microscope and the HORIBA LabRam HR800 Raman microspectrometer, respectively at the Institute of Geology and Geophysics, Chinese Academy of Sciences (IGGCAS). The cooling/heating stage was calibrated against the triple-point of pure CO₂ (-56.6 °C), the freezing point of water (0 °C) and the critical point of water (374.1 °C) using fluid inclusions synthesized by the FLUID Inc. Cooling/heating rates of 1 to 2 °C/min were adopted until phase transition when the rates were reduced to 0.1 - 0.2 °C/min. Carbonic phase melting (T_m, CO₂) and clathrate melting (T_m, clath) were determined by temperature cycling [59]. The precision of measurements was ± 0.2 °C at temperatures below 30 °C and ± 2 °C at temperature of the final homogenization. Fluid salinity and density were calculated based on T_m-ice and T_m-clath by using the Flincor program [60]. For the Raman analysis, an argon laser with a wavelength of 514.5 nm at power of 1000 mW were used. The spectral range of the composition analysis of fluid inclusions falls between 100 and 4000 cm⁻¹.

SEM-CL observations were undertaken at the Institute of Geochemistry, Chinese Academy of Sciences (IGCAS). The SEM was a JSM-7800F scanning electron microscope equipped with a Gatan Mono CL4 spectrometer and EDAX TEAM Apollo XL energy spectroscope. The carbo-coated and polished samples were analyzed at 10kV with beam current at 10nA, and the detection range of wavelength is from 300 to 700nm.

3.2. Quartz solubility diagrams

Diagrams of quartz solubility in this study are constructed for the H₂O-NaCl-CO₂ ternary systems. The temperature, pressure, composition (T-P-x_i) ranges of 100 to 600 °C, 1 to 4000 bar, salinity of 5 wt % NaCl, and 10 - 15 mol % CO₂ are selected based on previous [61,26,27] and the present fluid inclusion studies on the Linglong goldfield.

The phase boundaries between single phase liquid and coexistence of liquid and vapor in the H₂O-NaCl-CO₂ system were derived from equations of Mao et al. [62], which was based on the gas-liquid phase equilibrium of CO₂ in aqueous NaCl solutions. As the Linglong gold deposit has a relatively low salinity of an average value of 5 wt % NaCl, in the selected P-T range no NaCl saturation would occur.

Quartz solubility was calculated by using the density model of Akinfiev and Diamond [63] (AK model), which expressed as Eq. 1. Selection of this model was based on its simplicity, high accuracy and predictability. Despite several recent density models have been proposed [64,65], there are some inherent difficulties for application in salt- and gas-bearing solutions. The AK model has been successfully applied in several studies [11,18] for establishing quartz solubility and deposition framework for porphyry Cu deposits. In AK model, the temperature (T) and the molar volume of the mixed H₂O-NaCl-CO₂ hydrothermal fluid (V_{mix}) are the only two parameters for quartz solubility computing. The parameter V_{mix} of this study was acquired by using the equation of

state of Mao et al. [66,67]. V_{mix} can be parameterized with temperatures (T), pressures (P), fluid compositions (x_i), which means that the parameters of AK model are only T , P and x_i . Quartz solubilities are represented as isopleths in the corresponding diagrams, which were obtained by the following two steps: first, transform the AK model with the implicit function of T , P , x_i and quartz solubility (m_{SiO_2}); second, solve each specific quartz solubility with T - P relationships at the fixed fluid compositions with Newton Iteration.

$$\log m_{SiO_2} = A(T) + B(T) \times \log(18.0152/V_{mix}) + 2 \log x_{H_2O} \quad (1)$$

4. Results

4.1. CL textures and quartz sequences

Vein quartz in the four types of veins are examined by SEM-CL imaging, and the microtextures are used as petrographic guidance for constructing mineral paragenesis and quartz sequence, and for conducting fluid inclusion microthermometry. These textures are presented in Figure 5 and are summarized in the Table 2.

Based on CL luminosity in seventy-nine areas in twenty vein samples, two primary and three secondary textures in quartz have been identified. The first primary texture (PT1) is subhedral to euhedral concentric zonation of bright and dark bands, and quartz with this feature is categorized as Qz1 (Figure 5a, k); the second primary texture (PT2) is subhedral, dark CL, and quartz with this feature is designated as Qz2. The primary texture in Qz1 can be strongly obliterated where it is intensively fractured and dissolved by Qz2 (Figure 5a). Two subtypes for Qz2 can be distinguished according to occurrence and sulfide association. The Qz2a is represented by the quartz grains filled in relatively wider fractures (ST2), and in places caused brecciation of Qz1 (Figure 5c). They commonly intergrow with pyrite, gold and minor amounts of other sulfides (Figure 5b). Qz2b are those quartz grains fill in the voids networks in Qz1 (ST1) (Figure 5a, c), which are quite similar to the so-called "splatters" and "cobwebs" in early quartz grains in porphyry Cu deposits [8]. Another form of Qz2b is the pseudomorphic replacement of Qz1 (ST2, Figure 5e, g). These Qz2b quartz might have formed slight later than the "splatters and cobweb"-filling Qz2b, but no crosscutting relation are observed. But a difference is that the replacement Qz2b commonly show close spatial association with polymetallic sulfides (pyrite, sphalerite, galena, chalcopyrite, native gold and electrum), whereas the same association is absent for "splatters and cobweb"-filling Qz2b (Figure 5h, g).

In space, Qz1 primarily occurs in V1 quartz and V2 quartz-pyrite veins, but also occurs in low abundance in V3 quartz - polymetallic sulfide veins. Qz2a grains occur in variable amounts in all vein types, and are strongly associated with Qz1. Qz2a grains occur also in V1 and V2, but rarely in V3; In contrast, Qz2b grains occur primarily in V3 but rarely in V1 and V2.

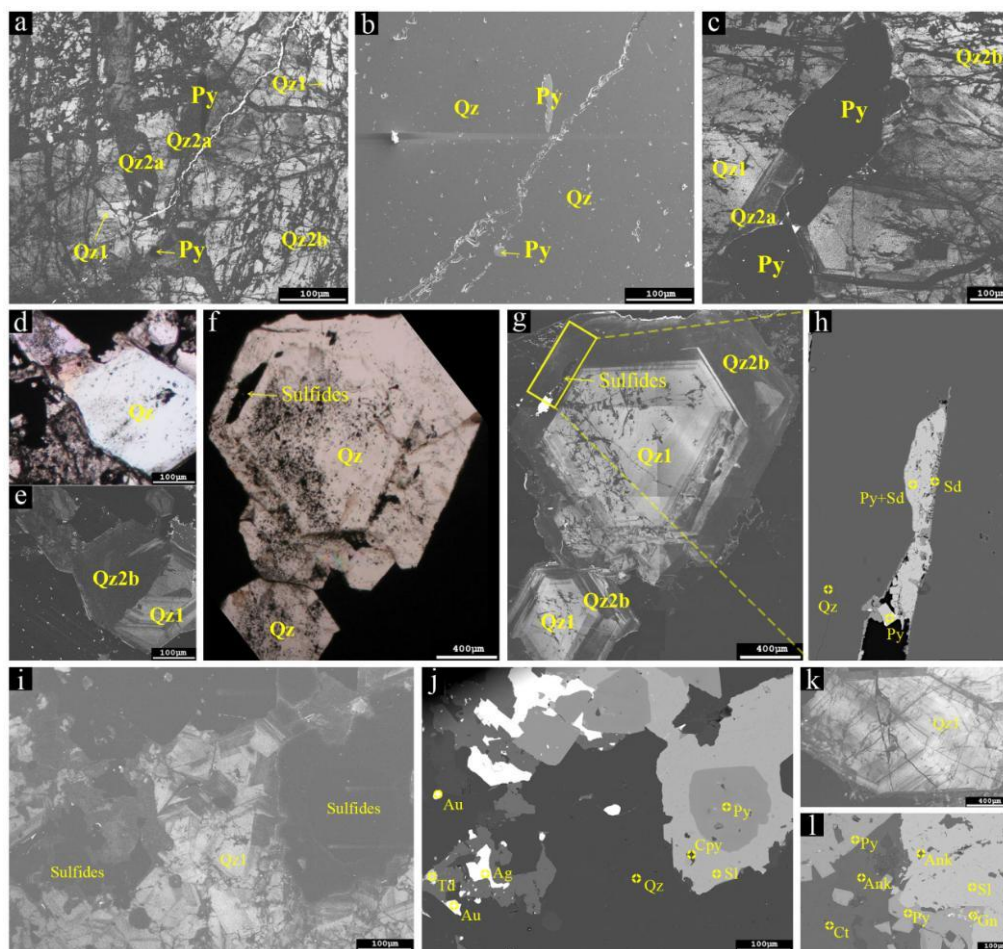


Figure 5. Microscopic and SEM-CL images of the quartz samples of the Linglong goldfield. (a) CL-bright Qz1 crosscut by later CL-gray to dark Qz2a in V1 vein. (b) secondary electron image of (a); (c) Pyrite grain intergrown with Qz2a in a V2 vein; (d) Euhedral quartz under transmitted light in a V2 vein; (e) Qz1 recrystallized by Qz2b with the same view as (d); (f) Euhedral quartz under transmitted light in V3 vein; (g) Qz1 recrystallized by Qz2b with the same view as image (f); (h) Backscattered electron (BSE) image with energy dispersive spectrometry (EDS) analyses of the area enlarged from (f); (i) Qz2b intergrown with sulfides in a V3 vein, with abundant relict Qz1; (j) EDS analysis of the same area of (i); (k) Qz1 in a V4 vein; (l) EDS analysis of minerals in a V4 vein. Quartz generations are detailed in the text. Py: pyrite, Sd: siderite, Cpy: chalcopyrite, Gn: galena, Sl: sphalerite, Td: tetrahedrite, Ag: Argentum, Au: Aurum, Ct: calcite; Ank: ankerite.

Table 2. Characteristics of vein types of the Linglong gold deposit.

	Quartz microscopic features	Quartz CL features	Associated sulfides	Associated alterations
V1	Euhedral-subhedral coarse grained; milky white and purify; fragmented	Bright (Qz1); Dark (Qz2a) Dark (Qz2b)	Pyrite	Potassic
V2	euhedral or subhedral grains; stress; white gray and poor transparency	Bright (Qz1); Dark (Qz2a)	Pyrite	Potassic ± sericitic ± pyritic ± silicic
V3	Subhedral grains; dark gray and content reduced	Bright (Qz1); Dark (Qz2a) Dark (Qz2b)	Pyrite ± chalcopyrite ± galena ± sphalerite ± magnetite ± hematite ± pyrrhotite	Potassic ± sericitic ± pyritic ± silicic
V4	White quartz	Bright, zoning	Pyrite (rare)	Carbonate

4.2. Fluid inclusion petrography and microthermometry

Based on

According to optical properties (phase species and ratios at room conditions), four types of fluid inclusions have been distinguished in the Linglong goldfield, i.e. pure CO₂ (A type), H₂O-NaCl-CO₂ liquid-vapor (B type), H₂O-NaCl liquid-vapor (C type), and daughter mineral-bearing liquid-vapor (D type) inclusions (Figure 6), which is consistent with previous studies [26,27]. The A type inclusions consist of monophase CO₂ (vapor or liquid) (Figure 6a), and two-phase CO₂ (L_{CO2}+G_{CO2}) (Figure 6b). This kind of inclusions are typically scattered in V3 and V4 with sizes of 5 - 10 μm. B type inclusions commonly consist of a carbonic phase and aqueous phase. According to the species of carbonic phase, B type inclusions are divided into B1 (L_{CO2}+G_{CO2}+L_{H2O}) (Figure 6c) and B2 (G_{CO2}+L_{H2O}) inclusions. B1 inclusions normally have volumetric ratio of carbonic phase 30 - 40% (Figure 6c, Figure 7a) whereas the ratios for B2 inclusions vary between 15 - 70% (Figure 6d, Figure 7a). B type of inclusions widely distributed in the whole vein generations and are relatively have a big size ranging from 10 - 20 μm. The presence of CO₂ phases is confirmed by Raman spectra (Figure 8a), and no other gas species are detected, which is consistent with microthermometric measurements. C type inclusions are normally consisted of a vapor phase along with a liquid phase, with variable V: L ratios and are developed from V2 to V4 with a middle size varied from 4 - 16 μm (Figure 6e). They can be subdivided into C1 (V: L > 50%) and C2 (V: L < 50% and variable) (Figure 7a). No CO₂ components have been detected for these inclusions by Raman spectrometry. In D type inclusions only appeared in V1 and V3 generations, besides a vapor and liquid phase (V: L ratios of 30 %), there exist a solid phase, which is commonly calcite as verified by laser Raman spectra (Figure 8b). The spatial distribution of fluid inclusions in quartz is revealed by CL imaging. In Qz1, B1 type inclusions occur along growth zones and are interpreted as primary in origin, and A and C types inclusions are also observed on or near healed cracks (Figure 7a), which are interpreted as secondary. Due to a scarce of inclusion in the studied Qz2a, no fluid inclusions can be distinguished as primary. In this case fluid inclusions observed in the V2 veins are taken for microthermometry, but cautions were taken when interpreting these data. In Qz2b, typical boiling assemblages were observed, with end-members of A type and C1 type and mixture of B2 type inclusions (Figure 7). Abundant C2 types are also present but maybe represent later fluids as supported by their low homogenization temperatures (Figure 7b).

During cooling and heating runs, A type inclusions showed carbonic phase melting at around -56.6 °C, indicating almost pure CO₂; clathrates were seldomly observed, indicating low concentrations of water; and the inclusion were finally homogenized between 27 - 30.9 °C; B1 type inclusions showed carbonic phase melting at -56.6 - -58.3 °C; clathrate melting occurred at 2.7 - 8.9 °C, corresponding to 2.8 - 12.3 wt % NaCl and final homogenization temperatures occurred at 260 - 330 °C; B2 type inclusions showed carbonic phase melting at -56.6 - -58.7 °C; clathrate melting at 2.8 - 10.0 °C, corresponding to 0 - 12.2 wt % NaCl and final homogenization occurred at 202 - 340 °C; C1 type inclusions showed ice melting at -5.6 - -4.6 °C, corresponding to 7.3 - 8.7 wt%NaCl, and final homogenization occurred at around 270 °C; C2 type inclusions showed ice melting at -4.8 - 0 °C, corresponding to 0.2 - 11.7 wt%NaCl, and final homogenization occurred at 120 - 280 °C; D type inclusions showed clathrate melting occurred at 2.7 - 7.0 °C, corresponding to 2.8 - 12.3 wt % NaCl, corresponding to 3.0 wt % NaCl, and finally homogenized at 210 - 315 °C. The microthermometric results of all fluid inclusions are listed the Table 3, but only those with CL image are plotted in the Figure 7b.

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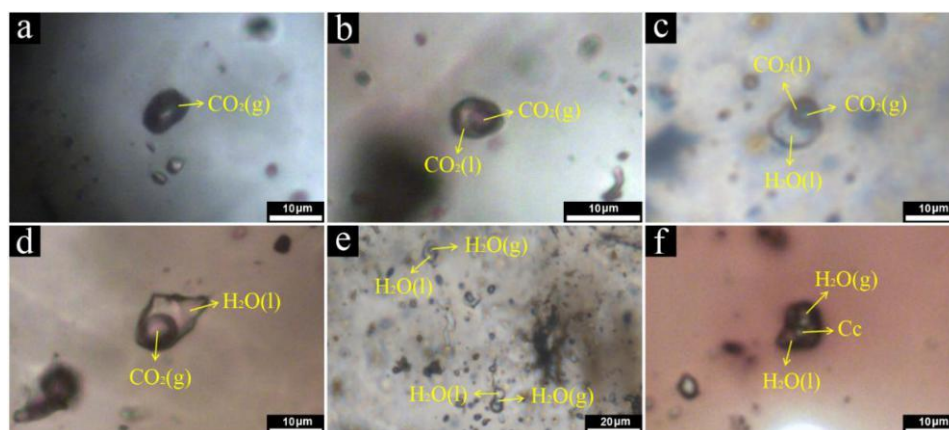


Figure 6. Microphotographs of different types of fluid inclusion at the Linglong gold deposit. A-type fluid inclusions with monophase CO₂ (Gco₂) (a) and two phases (Lco₂+Gco₂) (b); (c) B-type fluid inclusions with three phases (B1-type) (c) and two phases (B2-type) (d); (e) C1 (the upper) and C2 (the lower) fluid inclusions; (f) D type fluid inclusions. Cc: calcite.

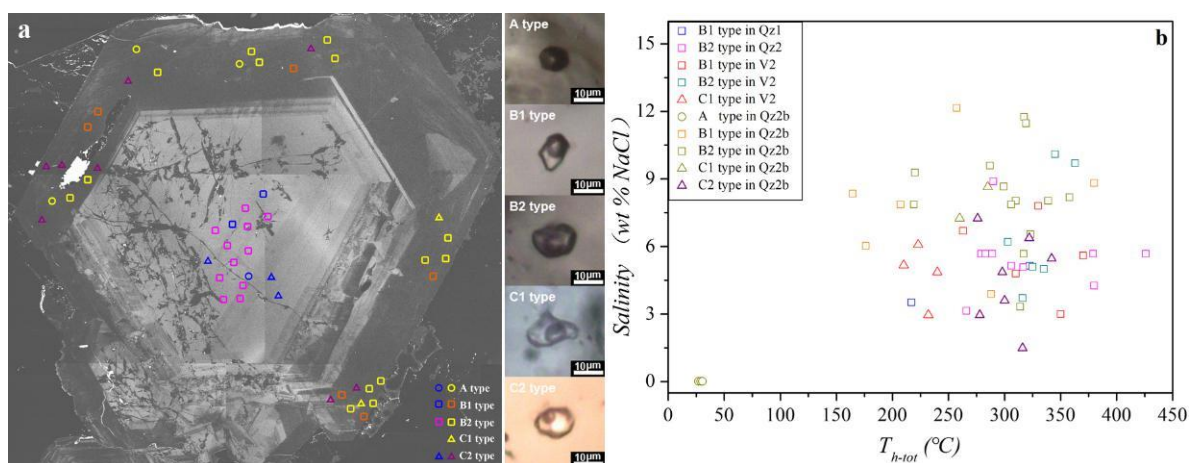


Figure 7. SEM-CL imaging of quartz with distributions of fluid inclusions (a) and a plot of homogenization temperatures and the salinity of different types of inclusions revealed by CL image in (a). Due to a lack of Qz2a in (a), fluid inclusion data are from Qz2a bearing V2 veins.

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Table 3. Microthermometric results of fluid inclusions at Linglong gold deposit.

Sample No.	Vein	FI type	Inclusion number	T _{m-CO2} (°C)	T _{m-ice} (°C)	T _{m-clath} (°C)	T _{h-CO2} (°C)	T _{h-tot} (°C)	Salinity (wt %)	CO ₂ density (g/cm ³)	Bulk density (g/cm ³)
17LL21	V1	A,	1	-57.0	/	/	/	27.5 (G)	/	0.28	/
		B1,	4	-57.2~ -57.0	/	2.7 ~ 8.6	29.7 ~ 30.9 (L)	299.6 (av.) (L or G)	2.76 ~ 12.28	0.53 ~ 0.60	0.72 ~ 1.00
		B2,	10	-58.7~ -57.4	/	4.5 ~ 9.0	28.2 ~ 30.9 (L)	317 (av.) (L)	2.0 ~ 9.74	0.53 ~ 0.65	0.81 ~ 0.98
10LL04	V2	D	1	/	/	2.7	/	208.2 (L)	12.28	0.53	0.92
		B2,	1	-57.0	/	5.6	/	269.5 (L)	9.14	/	/
		C2	2	/	-1.5, -1.0	/	/	282.9 (av.) (L)	1.65, 2.47	/	0.75, 0.76
17LL01	V2	B1,	6	-57.4 ~ -57.1	/	4.2 ~ 8.5	26.95 ~ 30.9 (G)	329.5 (av.) (L or G)	2.92 ~ 5.60	0.35 ~ 0.71	0.79 ~ 0.90
		B2,	12	-57.5 ~ -57.0	/	3.6 ~ 8.1	27 (L)	331.5 (av.) (L or G)	2.96 ~ 10.19	0.25 ~ 0.68	0.79 ~ 0.86
		C2	6	/	-2.2 ~ 0.1	4.2 ~ 8.5	/	263.4 (av.) (L)	0.17 ~ 3.60	/	0.69 ~ 0.95
17LL05	V2	B1,	3	-56.5 ~ -56.4	/	6.0 ~ 6.9	26.3 ~ 30.8 (G)	320.6 (av.) (L)	6.02 ~ 7.54	/	0.73 ~ 0.94
		B2,	4	-57.1 ~ -56.3	/	4.5 ~ 9.5	30.8 (G)	341.5 (av.) (L)	5.50 ~ 7.91	0.27 ~ 0.39	0.74 ~ 0.82
		C2,	2	/	-4.0, -2.2	/	/	222.6, 235.9 (L)	3.60 ~ 6.37	/	0.85 ~ 0.96
17LL09	V2	D	1	/	/	7.0	/	315.0 (L)	5.68	/	/
		B2,	4	-57.8 ~ -57.6	/	4.0 ~ 8.3	25.3 ~ 29.0 (L)	321.7 (av.) (L)	3.33 ~ 10.48	0.63 ~ 0.71	0.92 ~ 0.96
		A,	3	-56.8 ~ -56.7	/	4.7 ~ 7.0	/	27.0 ~ 30.5 (G)	/	0.27 ~ 0.37	/
17LL13	V3	B1,	8	-58.3 ~ -56.7	/	2.8 ~ 8.9	24.9 ~ 30.9 (G)	263.1(av.) (L or G)	3.89 ~ 8.35	0.53 ~ 0.76	0.74 ~ 0.94
		B2,	81	-57.7 ~ -56.7	/	2.8 ~ 10.0	/	287.1 (av.) (L or G)	0.02 ~ 12.15	0.53 ~ 0.64	0.87 ~ 0.95
		C1	2	/	-5.6 ~ -4.6	/	/	270 (av.) (L)	7.25 ~ 8.65	/	0.97 ~ 0.98
17LL15	V4	C2,	18	/	-8.0 ~ -0.5	/	/	255.3 (av.) (L)	0.83 ~ 11.70	/	0.60 ~ 1.0
		D	1	/	/	/	/	220.0 (L)	/	/	/
		A,	2	-56.7 ~ -56.6	/	/	/	29.9, 30.9 (L)	/	0.60, 0.59	/
12LL11	V4	B2,	14	-56.7~ -56.6	/	6.7 ~ 8.1	29.7 ~ 30.8 (G)	201.7(av.) (L)	3.7 ~ 6.2	0.33 ~ 0.40	0.65 ~ 0.81
		C2	2	/	-8.0~ -7.1	/	/	115.6, 143.1 (L)	10.6 ~ 11.7	/	1.0 ~ 1.03
12LL11	V4	C2	7	/	-7.8 ~ -2.3	/	/	144.0 (av.)	3.76 ~ 11.46	/	0.89 ~ 1.04

Note “/” denotes no data available.

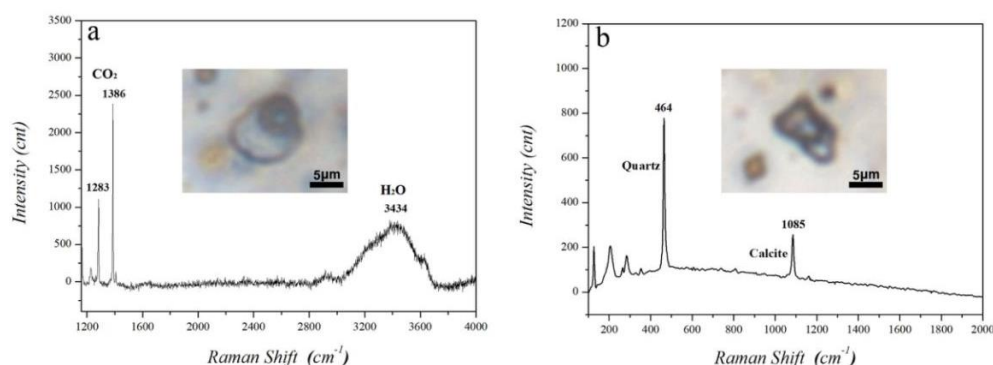


Figure 8. Raman spectra of B1 type fluid inclusion (a), and D type fluid inclusion fluid inclusions at the Linglong gold deposit (b).

4.3. Quartz solubility diagrams

Two types of solubility diagrams are constructed to study the effect of different parameters. P-T diagrams are constructed at constant fluid compositions (salinity: 5 wt % NaCl; X_{CO_2} : 0.1 and 0.15) to investigate effects of pressure and temperature (Figure 9). The second type is solubility versus temperature diagrams (Figure 10) at constant pressure to visualize effects of salinity and carbon dioxide contents.

In the P-T diagram (Figure 9), it is evident that solubility isopleths can be divided into three fields, i.e. T-sensitive yet P-inert field where isopleths are nearly vertical (TsPi); P-sensitive yet T-inert field where isopleths are nearly horizontal (PsTi); and retrograde solubility field where isopleths concave down towards (ReSo). The TsPi field occupies the high pressure region, and its boundary with the other two field shows linearly increase between P and T. The PsTi field occupies intermediate pressure and temperature region, while the ReSo field occupies low pressure region. It is noteworthy that the solubility isopleths show no discontinuity across the phase boundaries. In the TsPi field, the spacing between isopleths gradually increase with decreasing temperature. By contrast, spacing between the isopleths in PsTi field decreases with decreasing pressure, but at lower magnitudes than those for TsPi. It is noted that, in the studied temperature range, quartz solubilities are much lower at low pressures than those at high pressures. An increase of CO_2 molar content from 10% to 15% does not change the general topology, but shifts the temperature and pressure covers of the regions of TsPi, PsTi, and ReSo.

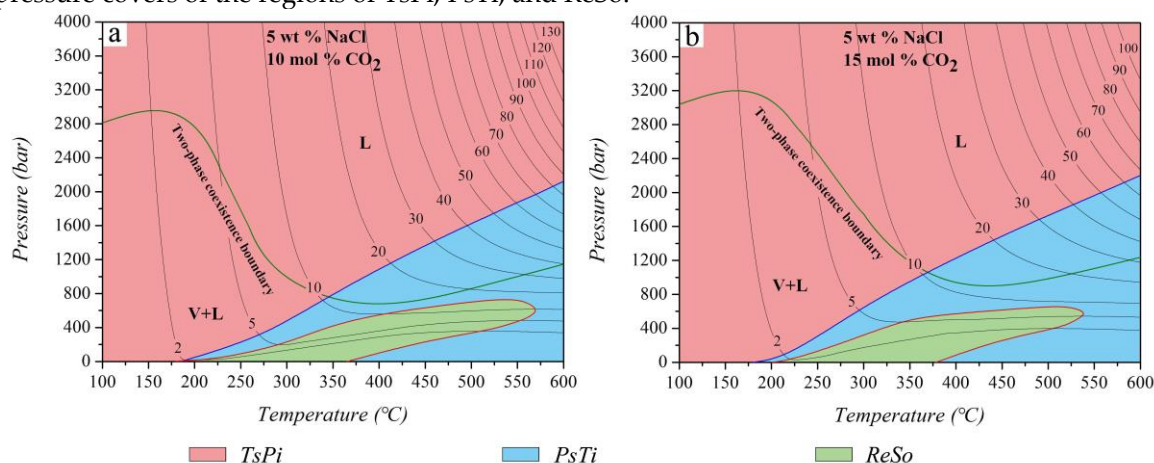


Figure 9. P-T diagram showing the single phase (liquid) and two-phase (vapor + liquid) field of H_2O -NaCl- CO_2 ternary system containing 10 mol % CO_2 (a) and 15 mol % CO_2 (b), contoured with quartz solubility isopleths. Three regions exhibiting different dependence on P and T are estimated (TsPi: temperature sensitive pressure inert; PsTi, and ReSo). For a fixed composition, if (T_0, P_0) is any point on the curve of equation (1), the boundary between TsPi and PsTi (the blue curve) is the locus

of points of isopleths that have $(\partial m / \partial T)_{P=P_0} = (\partial m / \partial P)_{T=T_0}$; the boundary between PsTi and ReSo (the red curve) is the locus of points of isopleths that have a slope of zero.

To investigate influence of X_{CO_2} and salinity, solubility versus temperature diagrams with three constant pressures (500 bar, 1000 bar, and 2000 bar), representing interception with different fields defined in the Figure 10, were constructed. CO_2 has negative effect on quartz solubility regardless of pressure and temperature (Figure 10a-c), due to its non-electrolyte nature that reduces silica hydration degree [63]. In addition, magnitude of solubility decrease among different X_{CO_2} dramatically reduces with decreasing temperatures, but increases with decreasing pressures. For instance, an increase of X_{CO_2} % from 5 % to 10% at 500 °C and 1000 bar can reduce quartz solubility by 10 mmol/kg (from 35 mmol/kg to 25 mmol/kg); whereas the magnitude in quartz solubility change drops to around 2.5 mmol/kg at 300 °C and 11 mmol/kg at 2000 bar. At 500 bars (Figure 10a), solutions with different X_{CO_2} % (between 0% mol and 15%mol) exhibit retrograde solubility between 600 °C and 350 °C. The magnitude of retrogression decreases with increasing X_{CO_2} . The peak point of quartz solubility shifts consistently to lower temperature with increasing X_{CO_2} . At 1000 bars and 2000 bars, no retrograde solubility is present, and instead solubility decreases monotonically (Figure 10b, c). It should be noted that the observed retrograde solubility patterns are a common feature for low pressures, not limited to 500 bars.

As to the effect of the electrolyte of NaCl, it demonstrates complex effect on quartz solubility by affecting the solution volume and changes hydration. At low pressure (500 bar) and the temperature approach the critical point of water, the quartz solubility demonstrates “salting-in” (quartz solubility enhanced when X_{NaCl} is increasing) effect and some extent of retrograde phenomenon (Fig 10d). At higher pressures, the “salting-in” effect of quartz solubility diminished progressively and the retrograde phenomenon is invisible (Fig 10e, f). As density models depict [63-65], at higher pressure conditions (> 2000 bars) the “salting-in” effect of quartz solubility will all changed into the “salting-out” (quartz solubility reduced when X_{NaCl} is increasing) effect because of the expanding of fluid molar volumes. Similar to CO_2 effect, magnitude of solubility decrease among different X_{NaCl} dramatically reduces with decreasing temperatures, but increases with decreasing pressures. However, the magnitude of solubility change is significantly lower than that of CO_2 effect (Figure 10d-f). Similarly, at low pressures (e.g. 500 bar), solution with 10%mol CO_2 and variable salinity displays retrograde solubility in the temperature range of 350 °C - 550 °C, and the peak point of solubility shifts to lower temperatures with decreasing salinity (Figure 10d). At higher pressures (e.g. 1000 bar and 2000 bar, Figure 10e, f), quartz solubility decreases monotonically.

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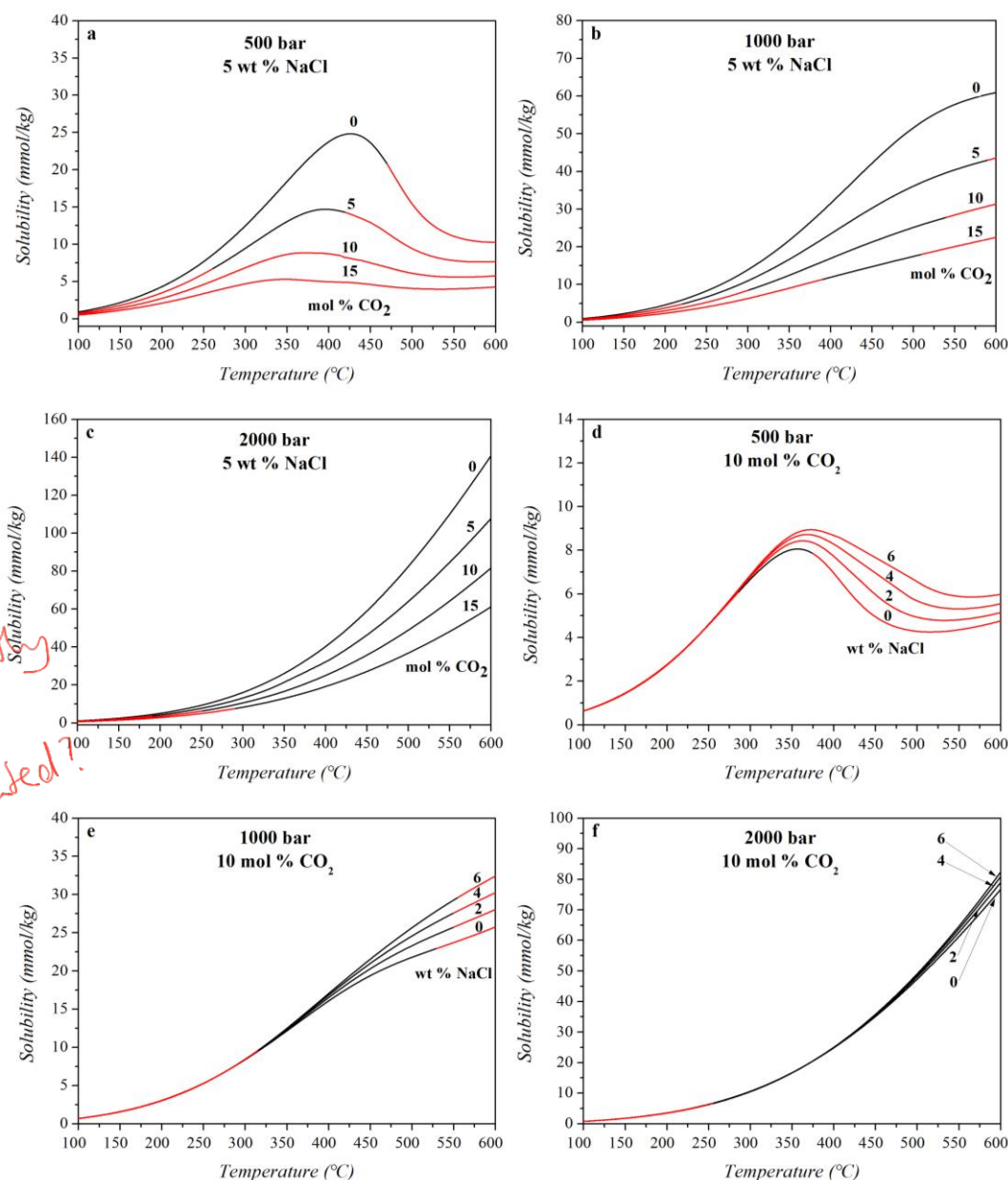


Figure 10. Predictions of quartz solubility in H₂O-CO₂-NaCl hydrothermal solutions. (a-c) are the fluids containing of 5 wt % NaCl and 0 - 15 mol % CO₂ compositions at 500, 1000, and 2000 bars conditions respectively; (d-f) are the fluids containing of 10 mol % CO₂ and 0 - 6 wt % NaCl compositions at 500, 1000, and 2000 bars conditions respectively; Notes: The computing results were obtained from AD model (2009). Black lines represent quartz solubility in liquid phase of hydrothermal fluids, and the red lines are quartz solubility in vapor + liquid phase of hydrothermal fluids. Quartz solubility in millimole SiO₂ per kilogram water.

5. Discussion

5.1. Fluid temperature and pressure estimated from fluid inclusions

Fluid inclusions have been used as a common means of estimating fluid pressure and temperatures. Normally, with phase-transition temperatures (e.g. T_h , $T_{m,ice}$) and appropriate PVTX diagrams or equations of state, minimum trapping temperature and pressure can be obtained. To know the exact trapping conditions, either independent P, T estimates such as those from mineral thermobarometers or estimates of rock load should be applied [68]. In the case of immiscible fluid

inclusion assemblage (coexistence of two fluids with contrasting density and composition), actual trapping conditions can be determined by intercept of isochores from the two types of inclusions.

In Qz1 of the Linglong goldfield, primary fluid inclusions are characterized by B-type liquid-vapor H_2O - NaCl - CO_2 inclusions (Figure 7a) that were trapped in single-phase field of a P-T diagram. Trapping conditions of this type of inclusion is estimated with complementary pressure information for the Jiaodong region. Liu et al. [45] studied the exhumation history of the Sanshandao gold deposit of the Jiaodong gold province since 110 Ma by an integrated method of structural analysis and (U-Th)/He thermochronology, and concluded that the deposit was eroded at least 5 km. A consistent result has been obtained by Li et al. [69], which suggested that the Daliuhang gold deposit may have exhumed 5 km since its formation by employing Al-in-hornblende barometer and geochronology. Using the intercepts between inclusion isochores of B1 inclusions in Qz1 and pressure (2.0 – 2.6 kbar) of Li et al. [69], the trapping temperatures are estimated to be 293 – 536 °C (the grey field in Figure 11a). This large temperature range might be attributed to uncertainty concerning if B1- or B2- type represents primary inclusions in Qz1. However, it is more likely that B1-type inclusions represent primary inclusions since B2-inclusions are commonly presented in Qz2b. Taking isochores of B1-type inclusions, the temperatures are narrowed to 388 – 502 °C (the dark field in Figure 11a).

In Qz2b, immiscible assemblage of fluid inclusions are present (coexistence of A and C1 inclusions), providing an opportunity to constrain the trapping conditions. Isochores of A and C1 inclusions are constructed, and their intercepts are taken as the estimates of trapping temperatures (292 – 331 °C) and pressures (273 – 431 bar) (Figure 11b). For Qz2a, due to a lack of CL-aided inclusion data, we tentatively infer that they were formed at similar temperatures but higher pressures (300 – 350 °C and 0.5 – 1.0 kbar).

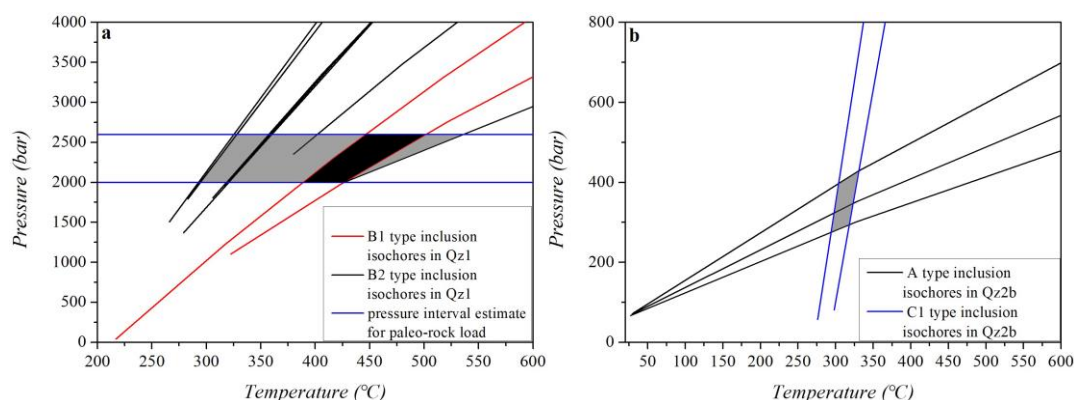


Figure 11. Estimated trapping pressures and temperatures by fluid inclusion isochores.

5.2. Quartz deposition and dissolution events at Linglong

Quartz deposition events at Linglong have been intensively studied and reported in previous studies [26,70], and a consistent conclusion has been reached. Three major quartz deposition events have been proposed based on macroscopic and microscopic observations of three generations of quartz-bearing veins, i.e. early barren quartz vein (V1 defined in this study), followed by pyrite quartz (V2) and polymetallic sulfide quartz (V3) veins. The same classification scheme has also been put forward for most gold deposits in the Jiaodong gold province [71]. A potential pitfall of using vein sequences to represent quartz sequence is that quartz of different generations might be not discernible by conventional observations [8]. With the aid of CL imaging, three generations of quartz are recognized for the Linglong deposit, which is consistent with the vein chronology. Detailed microtextural features are uncovered for each quartz deposition. The early quartz (Qz1) deposited during V1 quartz vein formation is characterized by relatively euhedral morphology (Figure 3a), bright CL and weak concentric zonings (Figure 12), likely suggesting deposition at relatively stable physico-chemical conditions (such as high temperature and slightly fluctuated

pressure, [72-74]). CL-aided fluid inclusion studies suggest that these quartz deposited from a single-phase carbonic aqueous solution at temperatures of 388 - 502 °C and pressures of 2.0 - 2.6 kbar. These quartz grains are commonly highly fractured and even brecciated, likely indicating that the early-formed V1 veins might have been reopened and served as conduits and depositional sites for later fluid events. During the course of Qz1 deposition, little metal sulfide and gold was precipitated.

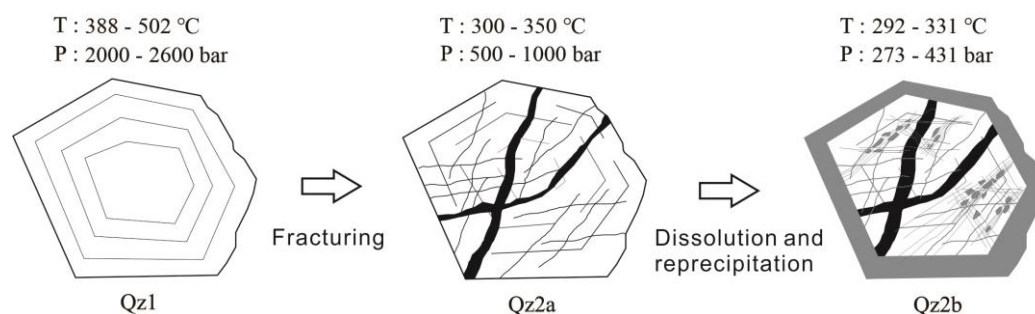


Figure 12. A sketch map of three different types of quartz in vein generations of Linglong goldfield.

The following quartz deposition events are represented by the formation of Qz2a and Qz2b. A significant difference revealed by CL study is that the amounts of these quartz are much more limited than that observed with conventional methods. More importantly, intimate and reliable spatial relations between these two quartz generations with metal sulfides and gold are identified. The dark CL of these quartz might indicate formation under conditions completely different from those of Qz1 (Figure 12), probably lower temperatures although evidence of trace elements are needed to make a quantitative conclusion [72-74]. Qz2a and Qz2b have no distinctive CL features except for the style of depositional site (Qz2a fills in brittle fractures, whereas Qz2b pseudomorphically replaces Qz1) and metal sulfide associations (Qz2a mainly intergrow with pyrite and possibly invisible gold, whereas Qz2b mainly coexist with polymetallic sulfides and visible gold). This might be an indication of slight yet significant changes in fluid properties and physico-chemical conditions. CL-aided fluid inclusion studies revealed that Qz2b was precipitated from immiscible fluids characterized by CO₂-rich vapor, H₂O-rich liquid and mixtures of the two at temperatures of 292 - 331 °C and pressures of 273 - 431 bar.

Besides quartz deposition events, a period of quartz dissolution is recognized through CL studies, which is the secondary CL textures of “splatters” and “cobweb” and pseudomorphic replacement in Qz1 (Figure 5a, c, e, g). The former texture in early quartz of porphyry Cu systems have been interpreted as a result of dissolution [8], while the latter might a result of combination of dissolution and reprecipitation. The dissolution event was immediately followed by deposition of Qz2b, which probably means that the dissolution process had created spaces for precipitation of quartz, sulfide, gold and other minerals.

5.3. Mechanism of early quartz deposition

CL-aided fluid inclusion petrography and microthermometry suggests that the early quartz (Qz1) was deposited at temperatures of 388 - 502 °C and pressures of 2.0 - 2.6 kbar with minor amounts of pyrite but little gold. The pressure temperature range indicates precipitation in the TsPi field of the Figure 9. In this field, pressure is a quite inefficient factor for changing quartz solubility. Salinity is also not efficient for controlling quartz solubility, especially at high pressures as demonstrated in Section 4.3 (Figure 10f). Therefore, among the parameters that may affect quartz solubility, temperature and CO₂ content may be decisive. Assuming the an original transporting fluid (10 mol % CO₂, 5 wt % NaCl) that carries SiO₂ at the solubility of 90 - 130 mmol/kg, at 550 - 600 °C and 3 - 3.5 kbar, decompressional cooling to the depositional conditions of Qz1 can

precipitate 70% to 80% silica, and this massive quartz precipitation can be further enhanced to over 90 wt % if CO₂ content is increased by 5 mol % (Figure 9, 13).

However, before attributing Qz1 quartz deposition to temperature, it is imperative to justify how cooling occurred and to what extent. Numerical considerations for ore formation have been undertaken by Toulmin and Mark [75] and Rose [76], whose conclusions may be used to interpret the case of Linglong. Rose [76] developed a model to investigate the effects of cooling 2.5×10^9 metric tons of solutions in a $0.5 \text{ km} \times 1 \text{ km} \times 1 \text{ km}$ rock volume at different fluid flow velocities from 700 to 400 °C for the formation of porphyry - style ore body. The conclusion reached was that simple cooling by conduction to wallrocks seemed not tenable. Instead, adiabatic cooling (the so-called throttling) and/or mixing with cool groundwater should be necessitated. This model works perfectly for the case of Linglong where the mass of solution involved (2.25×10^{10} metric tons, unpublished data), dimension of rock volume ($0.01 \text{ km} \times 1 \text{ km} \times 5 \text{ km}$) and temperature profile for an averaged quartz vein are comparable to those modeled parameters. Considering this, it is reckoned that cooling by conduction was not effective for Linglong, either. Moreover, adiabatic cooling is also considered impossible for Linglong since no evidence of dramatic pressure drop such as breccia has been observed. Similarly, mixing with cool groundwater is also considered unsuitable, because existing H-O isotopic studies consistently revealed dominance of magmatic components without meteoric involvement [26,27]. In addition, solubility models for pyrite and gold in H₂O- 10 wt % NaCl solution suggests that temperature has a strong negative effect on metal sulfides and gold (Figure 14a of Kouzmanov and Pokrovski, [77]), and thus cooling will inevitably lead to precipitation of sulfides and gold, which is contrary to the observations at Linglong. Therefore, the above discussions on thermal aspect and solubility behaviors of sulfides and gold may serve as arguments against the major role of cooling for the early veining at Linglong.

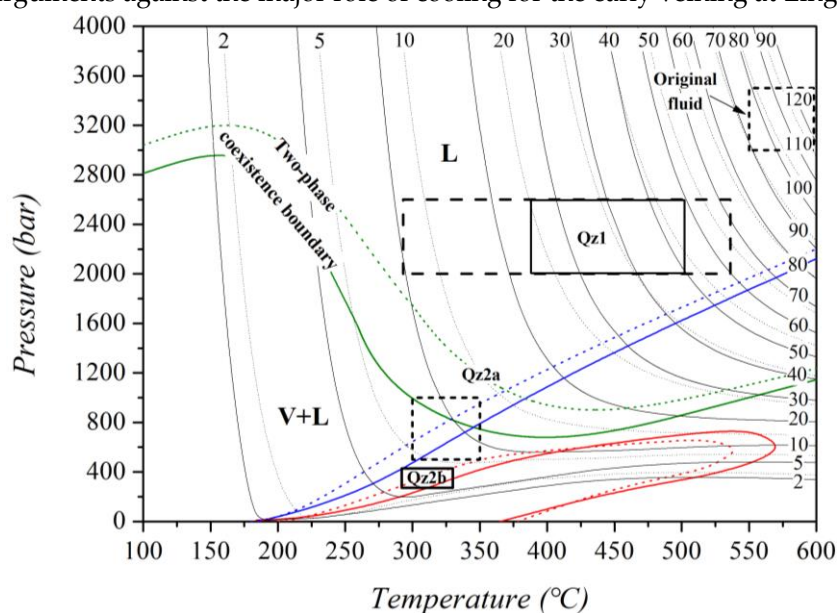


Figure 13. Formation of different veins in Linglong gold deposit. Notes: The black solid curves and dash curves represent quartz solubility isopleths in liquid and vapor + liquid hydrothermal fluids containing 10 mol % CO₂ and 5 wt % NaCl and 15 mol % CO₂ and 5 wt % NaCl solute, respectively. The green solid curves and dash curves, the blue solid curves and dash curves, and the red solid curves and dash curves are the boundaries of liquid phase and the vapor + liquid phase, of the TsPi and the PsTi, and of the PsTi and ReSo in different hydrothermal fluids containing 10 mol % CO₂ and 5 wt % NaCl and 15 mol % CO₂ and 5 wt % NaCl, respectively. Solid squares show the calculated conditions of Qz1 and Qz2b, while dashed squares show assumed conditions of original fluids, Qz1 and Qz2a.

Similarly, the possibility of CO₂ effect should also be evaluated. At Linglong, formation of Qz1 occurred along with alkaline metasomatism that significantly increased the amounts of secondary biotite. This formation of biotite may have preferentially consumed H₂O but exclude CO₂ and have

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led to increased CO₂/H₂O ratios, which is essentially CO₂ content in the fluid. As is shown in Figure 13, increase in CO₂ content will reduce quartz solubility, and therefore, it is likely that strong biotitization of wall rocks may have contributed to decrease of quartz solubility and thus quartz deposition. However, due to a lack of data on the scale of biotitization and thus it is impossible to quantitatively estimate CO₂ content change. If the acidity of the fluid system is internally buffered by CO₂ through the chemical reaction: $\text{CO}_2 + \text{H}_2\text{O} = \text{HCO}_3^- + \text{H}^+$ (2) as suggested by Phillips and Powell [78] and Phillips and Evans [79,80], biotitization will consume OH⁻ and thus reduce fluid pH value. The change in fluid acidity may explain the reversal behavior between quartz and pyrite / gold. As is shown by the solubility models of Kouzmanov and Pokrovski [77] (their Figure 14c), an reduce in pH will increase the solubility of pyrite and, to some extent, gold. Therefore, the biotitization (pH increase) may have been responsible for the deposition of Qz1 without pyrite and gold in the early veining process.

5.4. Mechanism of late quartz deposition

CL petrography has revealed two prominent features for Qz2a, i.e. occurrence in fractures that, in places, brecciated Qz1a; and significantly reduced amount in comparison to Qz1. The first feature indicate that the fluids may have experienced overpressure before the deposition of Qz2a. Fluid inclusion data of this study, although not as reliable as those for Qz1 and Qz2b obtained with the aid of CL, indicate that Qz2a was likely deposited at lower temperatures in the two-phase field and thus lowered quartz solubility (Figure 7). If the source remains the one as to that of Qz1, decompression alone would have caused quartz deposition in a scale similar to the early veining. It is plausible that the second feature argues for an either significantly cooled source or mixing with a silica - undersaturated fluid. The latter hypothesis is supported by H-O isotopic data, which revealed the presence of meteoric water during the formation of pyrite quartz veins (V2 of this study) [27]. Ingression of the silica - undersaturated meteoric water may have diluted the ore fluids and limited the amounts of precipitated quartz. Meanwhile, enhanced cooling induced by the meteoric ingression may have been responsible for the deposition of large quantities of sulfides and gold (Figure 14a of Kouzmanov and Pokrovski, [77]).

Because no distinctive difference in CL and chronological evidence are present for Qz2a and Qz2b, it is highly likely they represent a continuous fluid evolution. A marked feature of the Qz2b formation process is that it may have undergone ^{gone} quartz dissolution prior to reprecipitation as evidenced by the “splatter” and pseudomorphic replacement textures. According to the solubility behavior of quartz, dissolution would require: (1) heating in the TsPi region; (2) compression in the PsTi field; or (3) cooling through ReSo field. Fluid inclusion microthermometric data revealed that Qz2b might have been deposited at temperatures of 292 – 331 °C and pressures of 273 – 431 bar, corresponding to an overlapping region between PsTi and ReSo fields (Figure 13), ruling out the first hypothesis. However, no pressurization events have been suggested for Linglong, and thus the second hypothesis is considered impossible. Therefore, it is reasonable to infer that the quartz dissolution prior to Qz2b is considered as a result of cooling in the ReSo field of quartz solubility. In addition, CO₂ decrease may also be significant in increase quartz solubility. This process may have been due to vapor loss in the two-phase field (such as the model proposed by Hu et al. [71]) in the presence of fluid boiling in an open system, although the degree of this effect can not be evaluated due to a lack of relevant data. More importantly, the decrease in CO₂ may have been caused by phyllic alterations with consume H⁺ and thus move the Eq. (2) to the right side as opposed to biotitization. Qz2b, along with polymetallic sulfides and gold, deposition may have been induced by further cooling. To this end, a conclusion can be drawn that fluid cooling in ReSo field of quartz solubility combined with CO₂ loss might have been responsible for quartz dissolution, and continued cooling may have resulted in the late veining at Linglong.

6. Conclusion and Implications

Key findings of this study bear important implications for understanding genetic processes of gold veins in the Jiaodong gold province and other orogenic terrains at large. In Jiaodong, gold mineralization in most deposits, regardless of ore styles, occurred after a preceding major episode of Au-barren quartz deposition as thick quartz veins or as thin quartz veinlets [81]. The mechanism of this massive quartz depositional event and its role in the following deposition of metal sulfides and gold remains largely unattended. Quartz solubility interpretations coupled with CL-aided fluid inclusion analyses revealed an insignificant role of decompression due to “near vertical” isopleths and a limited role of fluid cooling regarding the efficiency of heat loss in the formation of this early quartz veining event. The exclusion of pressure, especially huge pressure drop induced by earthquakes [17], as a major factor for quartz deposition is based on quartz solubility analysis in carbonic brine solutions, which seems to be more accurate than those based on pure water. In the meantime, an increase in the content of CO₂ in fluid is highlighted due to its capacity of decreasing quartz solubility in the relevant physico-chemical conditions. This interpretation helps to establish a consequential link between quartz veining and the coeval biotitization that may have changed fluid pH, and provides a possible explanation for the Au-barren nature of this early veining event.

Another discovery made by this study is that the quantities of quartz that are genetically related to sulfides and gold is much lower than expected by most stable isotopic and fluid inclusion studies that have been based on conventional petrographic methods other than CL. This discovery should bring attention for future studies of gold deposits in orogenic intrusion-related settings to employ *in-situ* analytical techniques for obtaining stable isotopic and fluid inclusion data and to interpret the data on a firm basis of CL petrography. Such practice will significantly improve the quality of data and interpretations on fluid source and evolution, which are currently problematic due to fluid overprints [39]. In addition, CL microtextures suggest that the early quartz veins are very important site for later deposition of sulfides and gold, as evidenced by the fracturing, brecciation and replacement textures. In this sense, the early veining process are also important part of the entire fluid process and has a role to play for the gold genesis.

The roles of temperatures, pressures, phase separation, fluid-rock interaction and fluid mixing has been intensively discussed by previous studies of gold complexing, transport and deposition [19, 79,80]. This study offers a new perspective for addressing this issue with quartz solubility analysis. Detailed discussion on quartz dissolution and precipitation process for the main stages of gold formation indicates important roles of fluid mixing and fluid-rock interaction in the quartz veining process. These interpretations are aligned with previous studies aforementioned.

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References:

1. Götz, J. Classification, mineralogy and industrial potential of SiO₂ minerals and rocks. In Quartz: Deposits, Mineralogy and Analytics; Springer: Berlin, Heidelberg, 2012; pp. 1-27.

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2. Burley, S.D.; Mullis, J.T.; Matter, A. Timing diagenesis in the Tartan Reservoir (UK North Sea): constraints from combined cathodoluminescence microscopy and fluid inclusion studies. *Mar. Petro. Geol.* **1989**, *6*, 98-120.
3. Boiron, M.C.; Essarraj, S.; Sellier, E.; Cathelineau, M.; Lespinasse, M.; Poty, B. Identification of fluid inclusions in relation to their host microstructural domains in quartz by cathodoluminescence. *Geochim. Cosmochim. Ac.* **1992**, *56*, 175-185.
4. Götze J. Chemistry, textures and physical properties of quartz—geological interpretation and technical application. *Mineral Mag.* **2009**, *73*, 645–671.
5. Rusk, B. Cathodoluminescent textures and trace elements in hydrothermal quartz. In *Quartz: Deposits, mineralogy and analytics*, Springer, Berlin, Heidelberg, **2012**; pp. 307-329.
6. Pirajno, F. *Hydrothermal processes and mineral systems*; Springer Science & Business Media, 2009.
7. Van den Kerkhof, A.M.; Hein, U.F. Fluid inclusion petrography. *Lithos* **2001**, *55*, 27-47.
8. Rusk, B.; Reed, M. Scanning electron microscope-cathodoluminescence analysis of quartz reveals complex growth histories in veins from the Butte porphyry copper deposit, Montana. *Geology* **2002**, *30*, 727-730.
9. Frelinger, S.N.; Ledvina, M.D.; Kyle, J.R.; Zhao, D. Scanning electron microscopy cathodoluminescence of quartz: Principles, techniques and applications in ore geology. *Ore Geol. Rev.* **2015**, *65*, 840-852.
10. Fournier, R.O. Hydrothermal processes related to movement of fluid from plastic into brittle rock in the magmatic-epithermal environment. *Econ. Geol.* **1999**, *94*, 1193-1211.
11. Monecke, T.; Monecke, J.; Reynolds, T.J.; Tsuruoka, S.; Bennett, M.M.; Skewes, W.B.; Palin, R.M. Quartz Solubility in the H₂O-NaCl System: A Framework for Understanding Vein Formation in Porphyry Copper Deposits. *Econ. Geol.* **2018**, *113*, 1007-1046.
12. Maydagán, L.; Franchini, M.; Rusk, B.; Lentz, D.R.; McFarlane, C.; Impicini, A.; Ríos, F.J.; Rey, R. Porphyry to epithermal transition in the altar Cu-(Au-Mo) deposit, Argentina, studied by cathodoluminescence, LA-ICP-MS, and fluid inclusion analysis. *Econ. Geol.* **2015**, *110*, 889-923.
13. Mao, W.; Rusk, B.; Yang, F.; Zhang, M. Physical and chemical evolution of the Dabaoshan porphyry Mo Deposit, South China: Insights from fluid inclusions, cathodoluminescence, and trace elements in quartz. *Econ. Geol.* **2017**, *112*, 889-918.
14. Rusk, B.G.; Lowers, H.A.; Reed, M.H. Trace elements in hydrothermal quartz: Relationships to cathodoluminescent textures and insights into vein formation. *Geology* **2008**, *36*, 547-550.
15. Allan, M.M.; Yardley, B.W. Tracking meteoric infiltration into a magmatic-hydrothermal system: A cathodoluminescence, oxygen isotope and trace element study of quartz from Mt. Leyshon, Australia. *Chem. Geol.* **2007**, *240*, 343-360.
16. Li, Y.; Li, X.H.; Selby, D.; Li, J.W. Pulsed magmatic fluid release for the formation of porphyry deposits: Tracing fluid evolution in absolute time from the Tibetan Qulong Cu-Mo deposit. *Geology* **2018**, *46*, 7-10.
17. Weatherley, D.K.; Henley, R.W. Flash vaporization during earthquakes evidenced by gold deposits. *Nat. Geosci.* **2013**, *6*, 294–298.
18. Monecke, T.; Monecke, J.; Reynolds, T.J. The influence of CO₂ on the solubility of quartz in single-phase hydrothermal fluids: implications for the foemation of stockwork veins in porphyry copper deposits. *Econ. Geol.* **2019**, *114*, 1195-1206.
19. Pokrovski, G.S.; Akinfiev, N.N.; Borisova, A.Y.; Zotov, A.V.; Kouzmanov, K. Gold speciation and transport in geological fluids: Insights from experiments and physical-chemical modelling; *Geol. Soc.: London*, 2014; 9-70.
20. Chen, S.B.; Liu, Z.J.; Wang, Q.F.; Wang, F.J.; Xue, Y.S.; Xu, L.; Wang, J.P.; Zhu, B.L. Mineralization age geodynamic background for the Shangjiazhuang Mo deposit in the Jiaodong gold province, China. *Ore Geol. Rev.* **2017**, *80*, 876-890.
21. Sun, H.S.; Han, J.B.; Shen, Y.K.; Liu, L.; Leng, S.L.; Xu, C.; Yang, Q.M.; Ge, F.J.; Ouyang, S.B.; Deng X. Zircon (U-Th)/He age and its implication for post-mineralization exhumation degree of Linglong and Jiaojia goldfields, Northwest Jiaodong, China. *Earth Science* **2016**, *41*, 644-654. (in Chinese with English abstract)
22. Song, M.C.; Li, S.Z.; Santosh, M.; Zhao, S.J.; Yu, S. Yi, P.H.; Cui, S.X.; Lv, G.X.; Xun, J.X.; Song, Y.X.; Zhou, M.L. Types, characteristics and metallogensis of gold deposit in the Jiaodong peninsula, Eastern North China Craton. *Ore Geol. Rev.* **2015**, *65*, 612-625.

23. Li, L.; Li, S.R.; Santosh, M.; Zhu, J.; Suo, X.J. Early Jurassic decratonic gold metallogenesis in the eastern North China Craton: Constrains from S-Pb-C-D-O isotope systematic and pyrite Rb-Sr geochronology of the Guilaizhuang Te-Au deposit. *Ore Geol. Rev.* **2018**, *92*, 558-568.
24. Liu, W.M.; Wu, J.C.; Guo, W.C.; Xu, C.S.; Yu, Q.J.; Song, X.Y.; Li, J. The research of geological characteristics mineralization prospecting application in Dongshan deposit, Linglong gold field. *Gold Sci. Tech.* **2008**, *16*, 41-44. (in Chinese with English abstract)
25. Lin, J.T. Geological characteristics and deep prospecting potential analysis of Linglong gold deposit in Shandong. *Prospect. Tech.* **2018**, *3*, 107-108. (in Chinese with English abstract)
26. Wen, B.J.; Fan, H.R.; Santosh, M.; Hu, F.F.; Pirajno, F.; Yang, K.F. Genesis of two different types of gold mineralization in the Linglong gold field, China: constrains from geology, fluid inclusions and stable isotope. *Ore Geol. Rev.* **2015**, *65*, 643-658.
27. Guo, L.N.; Goldfarb, R.J.; Wang, Z.L.; Li, R.H.; Chen, B.H.; Li, J.L. A comparison of Jiaojia-and Linglong-type gold deposit ore-forming fluids: do they differ? *Ore Geol. Rev.* **2017**, *88*, 511-533.
28. Qiu, Y.S.; Yang, G.H.; Wang, K.H. Mineralization conditions of Zhaoyuan-Yexian Gold Belt in Shandong; Liaoning Science and Technology Publishing House: Shenyang, 1988; pp. 1-268.
29. Qiu, Y.M.; Groves, D.I.; McNaughton, N.J.; Wang, L.G.; Zhou, T.H. Nature, age, and tectonic setting of granitoid-hosted, orogenic gold deposits of the Jiaodong Peninsula, eastern North China Craton, China. *Mineral. Deposita* **2002**, *37*, 283-305.
30. Zhou, T.; Lü, G. Tectonics, granitoids and Mesozoic gold deposits in East Shandong, China. *Ore Geol. Rev.* **2000**, *16*, 71-90.
31. Fan, H.R.; Zhai, M.G.; Xie, Y.H.; Yang, J.H. Ore-forming fluids associated with granite-hosted gold mineralization at the Sanshandao deposit, Jiaodong gold province, China. *Mineral. Deposita* **2003**, *38*, 739-750.
32. Zhai, M.G.; Yang, J.H.; Liu, W.J. Large clusters of gold deposits and large-scale metallogenesis in the Jiaodong Peninsula, Eastern China. *Sci. China Ser. D: Earth Sciences* **2001**, *44*, 758-768.
33. Zhou, T.; Goldfarb, R.J.; Phillips, N.G. Tectonics and distribution of gold deposits in China—an overview. *Mineral. Deposita* **2002**, *37*, 249-282.
34. Zhou, X.H.; Yang, J.H.; Zhang, L.C. Metallogenesis of superlarge gold deposits in Jiaodong region and deep processes of subcontinental lithosphere beneath North China Craton in Mesozoic. *Sci. China Ser. D: Earth Sciences* **2003**, *46*, 14-25.
35. Yang, J.H.; Wu, F.Y.; Wilde, S.A. A review of the geodynamic setting of large-scale Late Mesozoic gold mineralization in the North China craton: an association with lithospheric thinning [J]. *Ore Geol. Rev.* **2003**, *23*, 125-152.
36. Mao, J.W.; Wang, Y.T.; Li, H.M.; Pirajno, F.; Zhang, Z.C.; Wang, R.T. The relationship of mantle-derived fluids to gold metallogenesis in the Jiaodong Peninsula: Evidence from D-O-C-S isotope systematics. *Ore Geol. Rev.* **2008**, *33*, 361-381.
37. Goldfarb, R.J.; Santosh, M. The dilemma of the Jiaodong gold deposits: are they unique? *Geosci. Front.* **2014**, *5*, 139-153.
38. Li, S.R.; Santosh, M. Metallogeny and craton destruction: records from the North China Craton. *Ore Geol. Rev.* **2014**, *56*, 376-414.
39. Goldfarb, R.J.; Groves, D.I. Orogenic gold: Common or evolving fluid and metal sources through time. *Lithos* **2015**, *233*, 2-26.
40. Groves, D.I.; Santosh, M. The giant Jiaodong gold province: the key to a unified model for orogenic gold deposits? *Geosci. Front.* **2016**, *7*, 409-417.
41. Deng, J.; Wang, C.; Bagas, L.; Santosh, M.; Yao, E. Crustal architecture and metallogenesis in the south-eastern North China Craton. *Earth-sci. Rev.* **2018**, *182*, 251-272.
42. Deng, J.; Wang, Q.; Santosh, M.; Liu, X.; Liang, Y.; Yang, L.; Yang, L. Remobilization of metasomatized mantle lithosphere: a new model for the Jiaodong gold province, eastern China. *Mineral. Deposita* **2019**, *1-18*.
43. Groves, D.I.; Santosh, M.; Deng, J.; Wang, Q.; Yang, L.; Zhang, L. A holistic model for the origin of orogenic gold deposits and its implications for exploration. *Mineral. Deposita* **2019a**, 1-18.
44. Groves, D.I.; Zhang, L.; Santosh, M. Subduction, mantle metasomatism, and gold: A dynamic and genetic conjunction. *Geol. Soc. Am. Bull.* **2019b**.

45. Liu, X.; Fan, H.R.; Evans, N.J.; Yang, K.F.; Danišić, M.; McInnes, B.I.; Qin, K.Z.; Yu, X.F. Exhumation history of the Sanshandao Au deposit, Jiaodong: constraints from structural analysis and (U-Th)/He thermochronology. *Sci. Rep.* **2017**, *7*, 1-12.
46. Goldfarb, R.J.; Hart, C.; Davis, G.; Groves, D. East Asian gold: deciphering the anomaly of Phanerozoic gold in Precambrian cratons. *Econ. Geol.* **2007**, *102*, 341-345.
47. Wu, F.Y.; Yang, J.H.; Xu, Y.G.; Wilde, S.A.; Walker, R.J. Destruction of the North China craton in the Mesozoic. *Annu. Rev. Earth Pl. Sci.* **2019**, *47*, 173-195.
48. Liu, F.L.; Liu, P.H.; Wang, F.; Liu, C.H.; Cai, J. Progresses and overviews of voluminous meta-sedimentary series within the Paleoproterozoic Jiao-Liao-Ji orogenic/mobile belt, North China Craton. *Acta Petrol. Sin.* **2015**, *31*, 2816-2846.
49. Zheng, J.P.; Sun, M.; Lu, F.X.; Pearson, N. Mesozoic lower crustal xenoliths and their significance in lithospheric evolution beneath the Sino-Korean Craton. *Tectonophysics* **2003**, *361*, 37-60.
50. Xie, S.W.; Wu, Y.B.; Zhang, Z.M.; Qin, Y.C.; Liu, X.C.; Wang, H.; Qin, Z.W.; Liu, Q.; Yang, S.H. U-Pb ages and trace elements of detrital zircons from Early Cretaceous sedimentary rocks in the Jiaolai Basin, north margin of the Sulu UHP terrain: provenances and tectonic implications. *Lithos* **2012**, *154*, 346-360.
51. Yang, K.F.; Fan, H.R.; Santosh, M.; Hu, F.F.; Wilde, S.A.; Lan, T.G.; Lu, L.N.; Liu, Y.S. Reactivation of the Archean lower crust: implications for zircon geochronology, elemental and Sr-Nd-Hf isotopic geochemistry of late Mesozoic granitoids from northwestern Jiaodong Terrane, the North China Craton. *Lithos* **2012**, *146*, 112-127.
52. Cai, Y.C.; Fan, H.R.; Santosh, M.; Liu, X.; Hu, F.F.; Yang, K.F.; Lan, T.G.; Yang, Y.H.; Liu, Y.S. Evolution of the lithospheric mantle beneath the southeastern North China Craton: constraints from mafic dikes in the Jiaobei terrain. *Gond. Res.* **2013**, *24*, 601-621.
53. Yang, L.Q.; Deng, J.; Wang, Z.L.; Guo, L.N.; Li, R.H.; Groves, D.I.; Danyushevskiy, L.; Zhang, C.; Zheng, X.L.; Zhao, H. Relationships between gold and pyrite at the Xincheng gold deposit, Jiaodong Peninsula, China: implications for gold source and deposition in a brittle epizonal environment. *Econ. Geol.* **2016**, *111*, 105-126.
54. Fan, H.R.; Hu, F.F.; Yang, J.H.; Zhai, M.G. Fluid evolution and large-scale gold metallogeny during Mesozoic tectonic transition in the Jiaodong Peninsula, eastern China; Geol. Soc.: London, Special Publications, 2007; 280, pp. 303-316.
55. Fan, H.R.; Feng, K.; Li, X.H.; Hu, F.F.; Yang, K.F. Mesozoic gold mineralization in the Jiaodong and Korean peninsulas. *Acta Petrol. Sin.* **2016**, *32*, 3225-3238.
56. Liu X.; Fan H.; Hu F.; Zheng X.; Lan T.; Yang K. Ore-forming fluid and stable isotope studies of Dazhuangzi gold deposit in Jiaodong Peninsula. *Mineral Deposits* **2011**, *30*, 675-715. (in Chinese with English abstract)
57. Li, R.H.; Albert, N.N.; Yun, M.H.; Meng, Y.S.; Du, H. Geological and Geochemical Characteristics of the Archean Basement-Hosted Gold Deposit in Pingliidian, Jiaodong Peninsula, Eastern China: Constraints on Auriferous Quartz-Vein Exploration. *Minerals* **2019**, *9*, 62.
58. Li, X.C.; Fan, H.R.; Santosh, M.; Hu, F.F.; Yang, K.F.; Lan, T.G. Hydrothermal alteration associated with Mesozoic granite-hosted gold mineralization at the Sanshandao deposit, Jiaodong Gold Province, China. *Ore Geol. Rev.* **2013**, *53*, 403-421.
59. Roedder, E., 1984. Fluid inclusions. *Rev. Mineral.* **12**, 1-664.
60. Brown, P.E.; Lamb, W.M. P-V-T properties of fluids in the system H₂O ± CO₂ ± NaCl: new graphic presentations and implications for fluid inclusion studies. *Geochim. Cosmochim. Acta* **1989**, *53*, 1209-1221.
61. Yang, Q.Y.; Shen, J.F.; Li, S.R.; Santosh, M.; Luo, Z.H.; Liu, Y. Oxygen, boron, chromium and niobium enrichment in native Au and Ag grains: A case study from the Linglong gold deposit, Jiaodong, eastern China. *J. Asian Earth Sci.* **2013**, *62*, 537-546.
62. Mao, S.D.; Zhang, D.H.; Li, Y.Q.; Liu, N.Q. An improved model for calculating CO₂ solubility in aqueous NaCl solutions and application to CO₂-H₂O-NaCl fluid inclusions. *Chemical Geology* **2013**, *347*, 43-58.
63. Akinfiev, N.N.; Diamond, L.W. A simple predictive model of quartz solubility in water-salt-CO₂ systems at temperatures up to 1000 °C and pressures up to 1000 MPa. *Geochim. Cosmochim. Acta* **2009**, *73*, 1597-1608.
64. Wei, Q.; Duan, Z.H.; Mao, S.D. A thermodynamic model of quartz solubility in H₂O-CO₂-NaCl systems up to 1000 °C and 1.5 GPa. *Acta Petrol. Sin.* **2012**, *28*, 2656-2666. (in Chinese with English abstract)

65. Shi, X.L.; Mao, S.D.; Hu, J.W.; Zhang, J.; Zheng, J.X. An accurate model for the solubilities of quartz in aqueous NaCl and/or CO₂ solutions at temperatures up to 1273 K and pressures up to 20,000 bar. *Chem. Geol.* **2019**, *513*, 73–87.
66. Mao, S.D.; Duan, Z.H.; Hu, J.W.; Zhang, D.H. A model for single-phase PVTx properties of CO₂–CH₄–C₂H₆–N₂–H₂O–NaCl fluid mixtures from 273 to 1273 K and from 1 to 5000 bar. *Chem. Geol.* **2010**, *275*, 148–160.
67. Mao, S.D.; Hu, J.W.; Zhang, Y.S.; Lv, M.X. A predictive model for the PVTx properties of CO₂–H₂O–NaCl fluid mixture up to high temperature and high pressure. *Appl. Geochem.* **2015**, *54*, 54–64.
68. Roedder, E.; Bodnar, R.J. Geologic pressure determinations from fluid inclusion studies. *Annu. Rev. Earth Pl. Sci.* **1980**, *8*, 263–301.
69. Li, X.H.; Fan, H.R.; Zhang, Y.W.; Hu, F.F.; Yang, K.F.; Liu, X.; Cai, Y.C.; Zhao, K.D. Rapid exhumation of the northern Jiaobei Terrane, North China Craton in the Early Cretaceous: insights from Al-in-hornblende barometry and U–Pb geochronology. *J. Asian Earth Sci.* **2018**, *160*, 365–379.
70. Zhang, Z.Q.; Lai, Y.; Chen, Y.J. Fluid inclusion study of the Linglong gold deposit, Shandong province, China. *Acta Petrol. Sin.* **2007**, *23*, 2207–2216. (in Chinese with English abstract)
71. Hu, H.L.; Fan, H.R.; Liu, X.; Cai, Y.C.; Yang, K.F.; Ma, W.D. Two-stage gold deposition in response to H₂S loss from a single fluid in the Sizhuang deposit (Jiaodong, China). *Ore Geol. Rev.* **2020**, 103450.
72. Wark, D.A.; Watson, E.B. TitaniQ: a titanium-in-quartz geothermometer. *Contrib. Mineral. Petro.* **2006**, *152*, 743–754.
73. Thomas, J.B.; Watson, E.B.; Spear, F.S.; Shemella, P.T.; Nayak, S.K.; Lanzirrotti, A. TitaniQ under pressure: the effect of pressure and temperature on the solubility of Ti in quartz. *Contrib. Mineral. Petro.* **2010**, *160*, 743–759.
74. Huang, R.; Audétat, A. The titanium-in-quartz (TitaniQ) thermobarometer: a critical examination and re-calibration. *Geochim. Cosmochim. Ac.* **2012**, *84*, 75–89.
75. Toulmin, P.; Clark, S.P. Thermal Aspects of Ore Formation, in *Geochemistry of Hydrothermal Ore Deposits*; Editor H. L. Barnes: New York, Holt, Rinehart and Winston, 1967; pp. 437–462.
76. Rose, A.W. Zonal relations of wallrock alteration and sulfide distribution at porphyry copper deposits. *Econ. Geol.* **1970**, *65*, 920–936.
77. Kouzmanov, K. and Pokrovski, G.S., 2012. Hydrothermal controls on metal distribution in porphyry Cu (–Mo–Au) systems. Special Publication of the Society of Economic Geologists, vol. 16, pp.573–618.
78. Phillips, G.N.; Powell, R. Link between gold provinces. *Econ. Geol.* **1993**, *88*, 1084–1098.
79. Phillips, G.N.; Evans, K.A. Role of CO₂ in the formation of gold deposits. *Nat.* **2004**, *429*, 860–863.
80. Phillips, G.N.; Powell, R. Formation of gold deposits: a metamorphic devolatilization model. *J. Metamorph. Geol.* **2010**, *28*, 689–718.
81. Fan, H.R.; Hu, F.F.; Yang, J.H.; Shen, K.; Zhai, M.G. Fluid evolution and large-scale gold metallogeny during Mesozoic tectonic transition in the eastern Shandong province. *Acta Petrol. Sin.* **2005**, *21*, 1317–1328. (in Chinese with English abstract)

